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RESPONSE OF AN INTERMOUNTAIN GROUNDWATER-DEPENDENT ECOSYSTEM TO WATER TABLE DRAWDOWN

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ABSTRACT.—Groundwater-dependent ecosystems (GDEs) that do not meet the legal definition of wetlands are important for sustaining regional biodiversity, livestock grazing, and outdoor recreation in the Intermountain West. Such GDEs in Owens Valley, California, are also used to produce 11,225 hectare meters (91,000 acre-feet) of water annually from about 100 water wells. We used 21 years of Landsat data and 18 years of field monitoring data to analyze responses of 2 adjacent-meadow GDEs to different groundwater management practices. The northern meadow, which was subject to continuous water table drawdown below the rooting zone of phreatophytic grasses, experienced decline in total live cover from 42.7% to 30.2%, decline in grass cover from 27.5% to 14.1%, transition from grass to shrub dominance, and change from groundwater dependence to precipitation dependence. These responses had been predicted by managers in 1976. The southern meadow, which was managed with cycles of water table drawdown and recovery, experienced neither cover decline nor dominance-type conversion and remained groundwater dependent. Variation in depth-to-water table (DTW) explained 83% of the pooled variance in total live cover in both meadows. Results showed that nonwetland, nonriparian GDEs are vulnerable to water table decline, as are wetland and riparian GDEs. Managing groundwater extraction through imposing one- to several-year cycles of water table drawdown and recovery may avoid further cover decline and type conversion in GDEs already affected by groundwater withdrawals.

RESUMEN.—En la región intermontañosa occidental, los ecosistemas dependientes del agua subterránea (de aquí en adelante denominados GDEs por sus siglas en inglés) que no cumplen con la definición legal de humedales son importantes para mantener la biodiversidad regional, el pastoreo de ganado y las actividades recreativas al aire libre. En Owens Valley, California, dichos GDEs también se utilizan para producir anualmente 112,250,000 metros cúbicos de agua de unos 100 pozos de agua. Utilizamos 21 años de datos Landsat y 18 años de datos de monitoreo en campo para analizar las respuestas de 2 prados adyacentes con GDEs a distintas prácticas de manejo de las aguas subterráneas. El prado del norte fue objeto de una reducción continua del nivel freático por debajo de la zona de enraizamiento de freatofitas. Este prado experimentó una disminución total de cobertura viva de un 42.7% a un 30.2%; la cubierta herbácea también disminuyó de un 27.5% a un 14.1%; a su vez, exhibió una tendencia que favorecía el dominio de arbustos y no de pastizales, y por último, cambió su dependencia de las aguas subterráneas a una dependencia de la precipitación. Estas respuestas habían sido pronosticadas por expertos en manejo en 1976. El prado del sur, el cual se controló mediante ciclos de reducción y recuperación de los niveles freáticos, no experimentó ni disminución de la cobertura ni ninguna clase de cambio con respecto al tipo de vegetación dominante y mantuvo su dependencia hacia las aguas subterráneas. En ambos prados, la variación en la profundidad de los niveles freáticos explicó el 84% de la varianza conjunta en el total de cobertura viva. Los resultados mostraron que los GDEs que no son humedales ni tampoco corresponden a vegetación ribereña son vulnerables al descenso de los niveles freáticos, tal como lo son los GDEs de las zonas ribereñas y de humedales. El control de la extracción de aguas subterráneas a través de la implementación de ciclos de reducción y recuperación de los niveles freáticos que duren de uno a varios años puede evitar que continúe la disminución de la cobertura y que cambie el tipo de vegetación dominante en los GDEs que están siendo afectados por la extracción de agua subterránea.

Not all groundwater-dependent ecosystems (GDEs) are equal in degree of groundwater dependence (Boulton and Hancock 2006); and GDEs such as wetlands and riparian zones, which are protected under the U.S. Clean Water Act, have been the subject of much ecological research (USEPA 2011). Nonwetland, nonriparian GDEs lack federal protection and have been studied less. Notwithstanding, such GDEs may be important for their biodiversity and economic value (City of Los Angeles and County of Inyo 1991a), especially in arid lands of the Intermountain West.

GDEs in Owens Valley, Inyo County, California, are a case in point. They cover over 23,000 ha, and most do not meet the legal wetland
definition (City of Los Angeles and County of Inyo 1991a). Groundwater is not expressed on the surface in these systems, and GDEs may be several kilometers from the nearest surface water. Following the classification of Eamus (2006), we will refer to these systems as Type III GDEs. Owens Valley Type III GDEs sustain a variety of sensitive species including endemics such as *Sidalcea covillei* (Owens Valley checkerbloom), *Calochortus excavatus* (Inyo County star-tulip), *Astragalus lentiginosus* var. *piscinensis* (Fish Slough milkvetch), and *Microtus californicus* ssp. *vallicola* (Owens Valley vole). Type III GDEs support livestock grazing and provide recreational opportunities (City of Los Angeles and County of Inyo 1991a).

Owens Valley Type III GDEs also encompass or adjoin nearly 100 wells, which are used to extract about 11,225 hectare meters (91,000 acre-feet) of water annually (Los Angeles Department of Water and Power 2011). Operation of these wells is conducted under the terms of the Inyo County/Los Angeles Long Term Water Agreement (Water Agreement)—a management agreement that allows water to be pumped, provided adverse impacts to GDEs are avoided (City of Los Angeles and County of Inyo 1991a). Management is based on a conceptual model of exploiting the ecological resilience of GDEs by imposing cycles of pumping-induced water table drawdown and recovery (Groeneveld 1992). Pursuant to the Water Agreement, certain sites have been the subject of intensive monitoring of vegetation and groundwater levels. In this paper, we use 21 years of Landsat and hydrologic data in conjunction with 18 years of field monitoring data to analyze vegetation response along a pumping-induced depth-to-water table (DTW) gradient. We chose our study site, which consists of 2 adjacent meadows, because it was dominated by native phreatophytic grasses when the DTW gradient was imposed and monitoring began (City of Los Angeles and County of Inyo 1991b). Changes in the meadows represent a case study in the responses of relatively intact (in regard to dominant species) Type III alkali-meadow GDEs to water table drawdowns of different magnitudes and durations. Our objectives were to document changes in vegetation cover and dominant life form in each meadow and analyze these changes in relation to DTW and precipitation. Results may be useful for GDE management in other arid regions.

**Study Site**

Owens Valley is a 120-km trough at the western edge of the Great Basin (Fig. 1). The Sierra Nevada forms the valley’s western wall and creates a rain shadow. Mean annual precipitation ranges from 127 to 152 mm, and more than three-quarters falls during the cold months, October–March (Hollett et al. 1991). The study site meadows lie at the toe of large, coalesced alluvial fans desceneding from the Sierran escarpment about 13 km north of the town of Independence, California. Substrate is a combination of ancient beach, bar, or river-channel sediments lying between alluvial fan deposits upslope to the west and fluvio-lacustrine deposits downslope to the east (Hollett et al. 1991). Soils are moderately to poorly drained loams (USDA–NRCS 2002), and elevation ranges from approximately 1178 m at the northwest corner to 1160 m at the southeast corner. The study site meadows cover about 204 ha and have been open to livestock grazing since Euro-American settlement in the 1860s. Livestock are removed in spring, approximately one month before vegetation data are collected. Analysis of postfire regrowth with grazing excluded (Pritchett and Manning 2009) and analysis of long-term monitoring data in grazing exclosures in both meadows (Inyo County Water Department 2011) suggest that the vegetation patterns and trends identified in our study occur independently of grazing effects.

Ecosystem conditions were first documented by C.H. Lee (1912). He noted the “striking correlation between vegetation and depth to groundwater,” and he mapped and described an extensive zone of grass-dominated meadow on the west side of Owens Valley. The meadow zone occurred on alluvial-fan, toe-slope areas where DTW was 2.4 m (8 ft) or shallower. Upslope, he documented a narrow zone of facultative phreatophytic shrubs, then Mojave Desert shrubland. All but the northwestern edge of our study site falls within Lee’s grass-dominated meadow zone (Fig. 1).

Lee (1912) also described the annual hydrographic signature of groundwater dependence. He observed in the meadow zone a water table decline of about 1 m commencing with the beginning of the growing season in spring, followed by rising water tables at the end of the growing season in autumn. Lee attributed
Fig. 1. Study site in central Owens Valley, California, including northern and southern meadow boundaries and monitoring wells used for kriging depth-to-water. Also shown is the 1912 area of shrub dominance and areas with depth-to-water >2.4 m—the value associated by Lee (1912) with the meadow–shrubland ecotone. Geographic coordinate system: Universal Transverse Mercator zone 11, NAD 27 datum.
this regular pattern of within-year water table drawdown and recovery to initiation and cessation of evapotranspiration. His data distinguished this signature of groundwater-dependence from deeper and irregular water table fluctuations upslope under precipitation-dependent shrub vegetation (Lee 1912).

The study site meadows were sampled by the Los Angeles Department of Water and Power in 1986 to characterize species composition and cover (Table 1). Dominant grasses were *Sporobolus airoides* (alkali sacaton) and *Distichlis spicata* (saltgrass) (nomenclature follows Baldwin et al. 2002, unless otherwise noted). The most common shrub species were *Atriplex lentiformis* ssp. *torreyi* (Nevada saltbush), *Ericameria nauseosa* (Pall. Ex. Pursh) G.L. Nesom & Baird (rubber rabbitbrush), and *Artemisia tridentata* Nutt. ssp. *tridentata* (basin big sagebrush). Less common species included *Glycyrrhiza lepidota* (American licorice), *Juncus balticus* (wire grass), and other herbaceous annual and perennial species (City of Los Angeles and County of Inyo 1991b). Pritchett and Manning (2009) presented a list of species recorded for the entire study site as of 2007. *Sporobolus airoides* and *D. spicata* are classified as "facultative" and "facultative wetland," respectively, with regard to wetland indicator status in California. "Facultative" is assigned to species equally likely to occur in wetland and nonwetlands, and "facultative wetland" is assigned to species usually occurring in wetlands. None of the shrubs are regarded as wetland indicators (USDA–NRCS 2011). The maximum effective rooting depths for Owens Valley phreatophytic shrubs and grasses were determined to be 3.7 m and 2 m, respectively (City of Los Angeles and County of Inyo 1991b).

Groundwater management of the study site meadows differed; however, the different treatments were a product of geography rather than management intention. The northern meadow is located southwest of 2 large pumps which operate continuously. The pumps are exempt from Water Agreement management protocols which would otherwise have required their operation to cease in order to facilitate water table recovery (City of Los Angeles and County of Inyo 1991b). High volumes of pumping during 1987–1989 from many wells in the area (Pritchett and Manning 2009) combined with drought conditions lowered the water table under the northern meadow to over 8 m below ground surface at rates of up to 110 cm \(\cdot\) year\(^{-1}\) (Fig. 2). With diminished pumping and the end of drought conditions, DTW recovered to about 5 m below ground surface. As a result, mean DTW under the northern meadow has remained well below the rooting zone of phreatophytic grass since 1988. In contrast, mean DTW under the southern meadow declined to only about 3 m at rates up to 70 cm \(\cdot\) year\(^{-1}\). During the 1990s, the water table rose to about 2 m; and since 1998, it has fluctuated near the bottom of the grass rooting zone (Fig. 2).

**METHODS**

Between 1984 and 1987, Los Angeles Department of Water and Power staff used air photos to circumscribe their Owens Valley land
into units of homogeneous vegetation referred to as “parcels.” The 2 study site meadows were defined as separate parcels in this project. Every parcel was sampled for species composition and top-layer canopy cover by running transects, 30.5 m (100 ft) in length, which were read using the point-intercept method (Bonham 1989). The data and parcel circumscriptions were subsequently accepted in the Water Agreement as management goals and baseline conditions against which pumping-induced changes would be measured. Although we follow Water Agreement usage and refer to these cover and composition data as baseline data, these data do not document undisturbed conditions. Vegetation in the study site meadows had already been subjected to intermittent periods of pumping-induced hydrologic alteration when baseline data were gathered (City of Los Angeles and County of Inyo 1991a).

Inyo County Water Department implemented the Water Agreement’s annual parcel monitoring protocol in 1991 to gather data for comparisons with conditions measured in baseline data. Inyo staff sampled top-level live canopy cover in selected parcels with randomly placed 50-m transects, which were read according to the point-intercept method (City of Los Angeles and County of Inyo 1991b). In both study site meadows, 12–26 transects were placed annually (1991–2007) in mid- to late June (Manning 2006). We refer to data from this sampling as “field cover.” The annual monitoring was designed to measure changes in dominant species. The methodology and relatively small number of transects did not

Fig. 2. Annual mean field total cover 1991–2007, 1986 baseline cover, pregrowing season precipitation, and mean depth-to-water for northern and southern study site meadows, Owens Valley, California. Error bars represent standard error.
reliably detect understory, smaller, and less abundant species. For these reasons we restricted response variables in our analyses to total live cover and cover of selected life forms. Total photosynthetic vegetation cover of both study site meadows was measured annually using linear spectral mixture analysis of Landsat Thematic Mapper data (Elmore et al. 2000). These data are referred to as “SMA cover.” Unlike field cover, they do not allow cover to be attributed to particular species or life forms. However, they have the advantage of being available for every year since the year baseline field data were gathered for the 2 study site meadows. Furthermore, Landsat data cover the full spatial extent of both meadows, thus representing a more complete sample than that obtained from the limited number of field transects. We used SMA cover from the period 1986–2006. Because Landsat data were collected in late August or September of each year and a fire burned large sections of both meadows in July 2007 (Pritchett and Manning 2009), we did not include 2007 SMA cover.

DTW and Precipitation Data

There are 14 groundwater monitoring wells in and adjoining the study site (Fig. 1), and we obtained their spring DTW values for the years 1986–2007 from Inyo County. Spring (ca. 1 April) values are well suited for inter-year studies of cover and DTW because they typically represent the annual water table high before initiation of within-year, evapotranspiration-induced drawdowns (Lee 1912). We used the spring DTW data as control points to create grids of kriged DTW estimates for the study area for each year, 1986–2007 (ESRI 2008). Kriging is recommended for interpolation in cases where sample sites are irregularly spaced (Legendre and Legendre 1998) and has been used by the Inyo County Water Department (Harrington and Howard 2000, Harrington 2003) to estimate mean spring DTW under vegetation parcels. We estimated mean DTW under each meadow by overlaying meadow boundaries on each annual grid of kriged DTW estimates and then calculating the mean of all grid cells in each meadow for each year.

We obtained daily precipitation data (cm), 1991–2007, from a Los Angeles Department of Water and Power rain gauge in the town of Independence, California, about 13 km to the south of the study site. We used water year precipitation (1 October–30 September) for correlations with SMA cover because SMA cover was gathered at the end of the water year. We used pregrowing season (1 October–30 June) precipitation for correlation with field cover (instead of water year precipitation) to exclude any precipitation which fell after sampling had been completed.

Data Analysis

COVER CHANGE.—We used transect data to calculate mean field total cover, grass cover (the sum of S. airoides and D. spicata), and shrub cover (all shrub species) for each study site meadow for each year, 1991–2007. Cover of other species was not included in this analysis; hence the sum of shrub cover and grass cover is slightly less than total cover. We regarded baseline means as standards without variance and applied one-sample t tests to test hypotheses that 1991–2007 mean field total, grass, and shrub covers were equal to their respective baseline values (Statsoft 2005). Baseline means were treated as standards without variance because transect locations were not random but had been selected as representative of the meadows (City of Los Angeles and County of Inyo 1991b). This limits their use for statistical inference (Zar 1984). Our use of these data is consistent with use by the Los Angeles Department of Water and Power (2011) and addresses its objections to the use of baseline and subsequent monitoring data (randomly sampled) in 2-sample t tests (Coufal 2001).

DOMINANCE-TYPE CONVERSION.—We divided grass cover by the sum of grass plus shrub cover to compute a grass proportion for each meadow and used results as an indicator of grass dominance. A value of 1 indicates all grass cover and no shrub cover; a value of 0 indicates no grass cover and all shrub cover. We used ordinary least-squares regression to examine the change in grass proportion over time (1991–2007) in each meadow. We computed the Durbin–Watson Statistic to test for the presence of serial correlation in the dependent variables in these regressions (Systat 2006).

COVER–WATER RELATIONSHIPS.—Due to the difficulty of normalizing distributions of precipitation and northern meadow DTW values, we used Spearman’s R to identify correlations of
field total cover, grass cover, and SMA cover with precipitation and with DTW in each meadow, 1991–2007 (1991–2006 for correlations involving SMA cover; Statsoft 2005). In addition to these individual meadow-based analyses of correlations, we pooled SMA cover data from both meadows to allow analysis of change in cover over the full range of DTW values in both meadows. We performed this analysis by means of ordinary least-squares regression (Systat 2006). The large range of DTW values made fitting data to a sigmoid curve more appropriate than fitting to a straight line (ter Braak 1996).

RESULTS

Cover Change

In the northern meadow, annual means of field total cover were below baseline in every year except 1998 (Fig. 2). Mean field total cover and mean grass cover over the entire monitoring period (1991–2007) were 30.2% and 14.1%, respectively. Both values were significantly below baseline values of 42.7% and 27.5%, respectively ($P < 0.01$). Mean field shrub cover over the same 1991–2007 period in the northern meadow was 11.3%, which did not differ significantly ($\alpha = 0.05$) from the baseline value of 10.4%. In the southern meadow, annual means of field total cover were at or above baseline in 11 of 17 years (Fig. 2). Mean field total cover, mean grass cover, and mean shrub cover, 1991–2007, were 53.4%, 41.1%, and 6.8%, respectively. None of these values differed significantly ($\alpha = 0.05$) from their respective baseline values (50.6%, 39.3%, and 6.1%). The 1991–2007 ranges of variation in annual mean field total cover were 42% and 41% for the northern and southern meadows, respectively (Fig. 2).

Dominance-Type Conversion

The meadows differed regarding trends in dominant life form. In the northern meadow, there was a significant negative trend in grass proportion over time ($R^2 = 0.72$, $P < 0.01$). There was not a significant trend ($P = 0.11$) in the southern meadow (Fig. 3). Values of the Durbin–Watson statistic were 2.21 and 1.52, respectively, showing little evidence of serial correlation (Systat 2006).

Cover–Water Relationships

The meadows differed regarding cover response to water. In the northern meadow, SMA, field total, and grass cover were significantly positively correlated ($P < 0.05$) with precipitation (Spearman’s $R = 0.66$, 0.63, and
0.55, respectively) but not with DTW. In the southern meadow, SMA, field total, and grass cover were significantly negatively correlated \((P < 0.05)\) with DTW (Spearman’s \(R = -0.81, -0.60,\) and \(-0.50,\) respectively) but not with precipitation. DTW explained 83% of the pooled SMA cover variation in both meadows over the entire 1986–2006 period (Fig. 4).

**DISCUSSION**

**Cover Change and Dominance-Type Conversion**

The magnitudes of annual cover fluctuations in our data are not uncommon for Owens Valley well field meadows, as measured with the relatively small sample sizes of the Water Agreement monitoring protocol (Manning 2006). More important is that 1991–2007 cover values in the southern meadow fluctuated around the 1986 management baseline value, while cover values in the northern meadow fluctuated around a mean about 13% lower than baseline. This result is consistent with photographic evidence (Inyo County Water Department 2011). Observations of Stromberg et al. (1996), Elmore et al. (2003), and Cooper et al. (2006) also associate water table declines with cover decline. Especially relevant were predictions made in 1976 by the Los Angeles Department of Water and Power for its Owens Valley holdings, including the study site meadows. The department defined general classes of predicted responses to pumping-induced water table declines and associated them with 3 ranges of DTW values (Los Angeles Department of Water and Power 1976). The magnitude and duration of water table drawdowns in the northern meadow placed it in the “extreme” category of predicted responses. Phreatophytic species, “would definitely thin out and begin to die. Evidence of this change would be expected within three to five years, with definite marked changes in 10 years [after water table drawdowns].” The magnitude and duration of drawdowns under the southern meadow placed it in the “none-moderate-high” category of predicted responses. The fluctuations of mean field total cover and mean grass cover around baseline were consistent with this prediction.

The decline of grass proportion over time in the northern meadow (Fig. 3) indicated conversion from meadow to shrubland, but no similar trend was seen in the southern meadow. Groeneveld (1992) warned that management under the Water Agreement may accelerate meadow conversion to shrubland because deep-rooted shrubs would be less vulnerable than
grasses to water table declines and would also be able to respond more quickly than grasses when water tables begin to recover. The conversion from grass to shrub dominance in the northern meadow was also consistent with observations of Lee (1912), who mapped shrublands in areas with DTW deeper than 2.4 m—a value he associated with the shrubland–meadow ecotone. Almost half the northern meadow had DTW >2.4 m in 1986 (Fig. 1). Instead of managing groundwater to return the 2.4 m DTW contour to its historic location at the northwest edge of the meadow, managers extracted groundwater (1986–2007) at rates which increased mean DTW under the entire northern meadow (and parts of the southern meadow) to levels exceeding 2.4 m. Groeneveld and Or (1994) associated an Owens Valley meadow–shrubland ecotone with a DTW threshold of approximately 1 m, even shallower than Lee’s.

West and Young (2000) cited a model of succession around retreating (drying) Great Basin lakes in which succulent and graminoid species give way to shrublands along water and salinity gradients. In the northern meadow, grasses *Distichlis spicata* and *Sporobolus airoides* declined in cover relative to shrubs *Atriplex lentiformis* ssp. *torreyi*, *Ericameria nauseosa*, and *Artemisia tridentata*. These are all taxa in the succession model cited by West and Young (2000). The Los Angeles Department of Water and Power (1976) also predicted “plant succession shifts toward non-phreatophytes” and a “definite shift toward xerism” as a response to water table decline. Consistent with this prediction, the encroaching shrub species in the northern meadow were more xerophytic (lacking wetland indicator status values), than the grasses they replaced. Type conversion in the northern meadow can be understood as a pumping-induced acceleration of succession, which formerly occurred as a response to long-term climatic drying. Froend and Sommer (2010) described response of an Australian GDE to water table decline up to 50 cm · year⁻¹ as a “transition to an alternative ecohydrological state.” The northern meadow was subject to drawdowns up to 110 cm · year⁻¹, and the resulting cover decline and type conversion makes “transition to an alternative ecohydrological state” an apt descriptor. Odion et al. (1992), Stromberg et al. (1996), Naumburg et al. (2005), Cooper et al. (2006), and Patten et al. (2008) all documented and/or predicted plant community change in response to long-term drawdowns.

**Cover–Water Relationships**

Although both study site meadows were classified as groundwater dependent under the Water Agreement on the basis of field data from 1986, the Spearman’s R values demonstrate that the northern meadow was statistically precipitation dependent during 1991–2007, while the southern meadow remained groundwater dependent. The analysis of pooled data integrated the annual SMA cover and DTW values of both meadows in a single model (Fig. 4) showing the shape of the cover–DTW response curve over the full range of DTW values typically imposed on Owens Valley well meadows (Manning 2006). The sigmoid curve describing this shape was also used by Shafroth et al. (2000) to describe the response of woody riparian vegetation to declining water table. The stem of the curve represents a range of relatively shallow DTW values where cover is correlated with changes in DTW. When DTW is in this range of values, groundwater is available to plant roots and overrides any effects of variation in precipitation on cover, as measured by methods used in this study. The tail of the sigmoid curve defines a range of DTW values where changes in DTW no longer affect cover because groundwater is too deep to be accessible to vegetation. In this range of values, effects of variation in precipitation on cover are measurable. Note that, under current climatic conditions, precipitation alone is not sufficient to sustain cover values as high as baseline values (City of Los Angeles and County of Inyo 1991b). As a result, when DTW is drawn down to the tail of the curve, it effectively limits cover, even as cover becomes correlated with precipitation.

Elmore et al. (2006) used SMA cover to identify 2.5 m as the mean threshold DTW value for meadow vegetation in Owens Valley above which there was measurable groundwater dependence and below which there was measurable precipitation dependence. Our results, using both field cover and SMA cover, are consistent with this threshold. Mean DTW under the northern meadow was close to the Elmore et al. (2006) range of groundwater dependence in 1986 and 1987, then mean DTW increased into the range of precipitation dependence for
the duration of the study period (Fig 4). Cover responded accordingly as the meadow converted from groundwater to precipitation dependence. The response had been anticipated by the Los Angeles Department of Water and Power (1976), which predicted that vegetation “will move toward equilibrium with the desert environment and prevailing 5-inch rainfall potential” in areas subject to drawdowns of the magnitude and duration of those imposed on the northern meadow.

Management Implications

When the Water Agreement was adopted, the conceptual model it applied to management of Type III GDEs had not been tested (City of Los Angeles and County of Inyo 1991a). Our analysis of southern meadow monitoring data demonstrated that the Water Agreement conceptual model of exploiting ecosystem resilience could meet management goals with regard to total cover, grass dominance, and groundwater dependence over a 21-year period. Our analysis of northern meadow monitoring data, where water tables were drawn down well below the grass rooting zone continuously for 19 years, showed that this area was unable to sustain management baselines regarding cover, dominance type, and groundwater dependence. Our analysis of pooled data showed the overall importance ($R^2 = 0.83$) of groundwater in both sustaining and limiting cover (relative to baseline values) depending upon whether groundwater is in the rooting zone or drawn down below it.

Though management goals were met in the southern meadow, the applicability of the Water Agreement conceptual model elsewhere has important limitations. The mid-1986 baselines for the study site meadows represented disturbed conditions. They were set after the meadows had already been subjected to 14 years of relatively intensive pumping. Furthermore, baseline samples were small (Table 1) and not intended to fully document the diversity within each meadow (City of Los Angeles and County of Inyo 1991b). As a result, management success in the southern meadow was due, in part, to the coarse scale of management goals. We could not recommend application of the Water Agreement conceptual model to GDEs not already subject to impacts of hydrologic alteration.

An important unanswered question concerns the duration necessary for recovery after a deep drawdown such as that experienced in the northern meadow. The Los Angeles Department of Water and Power (1976) predicted “substantially longer periods [>10 years] of normal conditions would be required for vegetation restoration.” The Inyo County Water Department and the Los Angeles Department of Water and Power have been attempting since 2007 to reach agreement on optimum durations of cycles of drawdown and recovery but have yet to do so. Had the partial groundwater recovery of 1998 been allowed to continue to the grass rooting zone, data relevant to this question would have been obtained.

Conclusion

Given the increasing pressure to exploit groundwater resources in the arid lands of western North America (SNWA 2011) and globally (Postel 2000), it is important to recognize that Type III GDEs are vulnerable to groundwater withdrawals just as their more widely studied wetland and riparian GDE relatives. Long-term data regarding Type III GDEs had been lacking (Groeneveld 1992), but the data presented here help fill this void. Our results show the response of groundwater-dependent alkali meadow to water table declines of different magnitudes and duration. As seen in the southern meadow, short-term cycles of water table drawdown and recovery may avoid or at least defer measurable changes in cover and dominance type in the meadow ecosystem. Prolonged water table drawdown under the northern meadow resulted in conversion to an altered ecohydrological state—a response consistent with management predictions (Los Angeles Department of Water and Power 1976) but inconsistent with management objectives (City of Los Angeles and County of Inyo 1991a).

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